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Project 3.2: Real-time Spatial Information Evaluation and Processing

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Outline

- Soil Moisture
 - Activities at the Kensaton soil moisture SMAP calibration/validation site
 - SMAP product validation over an Arctic tundra site
 - Cal/val field campaign in 2016
- Soil Freeze/Thaw
 - Ground-based networks
 - A soil freeze/haw validation campaign for SMAP





Kenaston Network

- SMAP consistently demonstrating a dry bias
- Strong performance at Kenaston network site: ubRMSE = 0.027, r = 0.79
- Global average ubRMSE is 0.038; r = 0.78 (meeting mission targets)
- SMAP 9km enhance product now available
- Entering senior review







Knowledge of the network CV could provide insight into time periods when the network should be used with caution for cal/val activities





The range in upscaled soil moisture across all possible random combinations of stations from 5 to 30. Soil weighting shows the least stability .





Table 1. The Pearson r for the Tb Provided on the Two Global Grid Pixels (G-1 and G-2) Associated With the TUP and TMM Sites, Respectively, and the North Polar Grid, Relative to the CMEM Modeled Tb

	Product	2015	2016
SMAP L1C Tb (TUP) versus CMEM	H-pol fore	0.02	0.10
	V-pol fore	-0.01	0.11
SMAP L1C Tb (TMM) versus CMEM	H-pol fore	0.26	0.11
	V-pol fore	0.18*	0.11
SMAP L1C Tb (North Polar grid) versus CMEM	H-pol fore	0.39*	0.51*
	V-pol fore	0.44*	0.55*
*Significance at α = 0.05.	Wrona	Wrona et al., 2017 GRL	
	W		Flood
			NSEF

Validation of the Soil Freeze Thaw State: Results



Air Temperature (top panel) good proxy under dry snow conditions, however very poor under wet or bare soil.

Under Bare soil conditions, soil dielectric measurements (near but not at surface) correspond most closely

Under dry snow conditions, soil temperature measurements near surface correspond most closely

Under wet snow both soil temperature and soil dielectric are similar, both greatly exceed air temperature proxy



Soil Freeze Thaw SMAP Experiment

- October 28-November 13, 2015 we conducted a freeze thaw SMAP experiment near Carmen, Manitoba
- Morning and afternoon flights of NASA's King-Air aircraft carrying their Scanning L-band active passive sensor (SLAP)
- Ground crews observing freeze/thaw state (temperature), soil freezing state and soil moisture









to diurnal F/T events for near surface freezing



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Characterizing Runoff Ratios over Canadian Watersheds using Satellite Derived Soil Moisture

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Soil Moisture-Runoff Linkages

- Soil moisture is widely recognized as one of the critical controls on surface runoff processes
- Influences the flashiness of streams within a catchment area
- Traditionally measured in situ (e.g. impedance probes)
- Soil-moisture assimilation into rainfall-runoff models have shown the value of incorporating soil moisture into hydrological floodforecast models

- Canada has a lack of real-time and historic observational runoff data over catchments
- Very sparse gauging network with complete records



FIGURE 11. ANNUAL UNIT RUNOFF (dam?/km?) FOR A 25% PROBABILITY OF EXCEEDENCE

Average Annual Unit Runoff dam³/km² Probability of exceedence - 1994 FloodNet

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SMOS Soil Moisture

Launched by the ESA in 2009

L-band radiometer operating at 1.4 GHz (21 cm wavelength)

Sun-synchronous, quasi-circular orbit, dusk-dawn, 23-day repeat cycle, **3-day sub-cycle**

Average spatial resolution of **40km**

Soil Moisture retrieval accuracy of 4%

Retrieves surface soil moisture content (top ~5cm of the soil)

Soil moisture signal impacted by snow, ice and frozen soils, vegetation cover.





Agriculture and Agri-food Canada, 2016; Ke r^{SeE} $a_{1,2}^{C}$ 2012

Research Gaps

- Soil moisture-runoff studies predominantly focused on a single instrumented basin
- Lack of data over large and remote regions of Canada

Research Aims

- Analyze the importance of SMOS antecedent soil moisture data in relation to runoff generation in a large scale Canadian context
- Characterize the dominant hydrologic and topographic controls on this relationship



Basin Selection



76 gauging stations from the Canadian Reference Hydrologic **Basin Network** (RHBN) and their contributing watersheds, distributed throughout the West Coast, Prairie, Central and Atlantic regions of Canada south of the 60th parallel. Watershed boundaries were determined from the HydroBASINS dataset (www.hydrosheds.org/)



Data Sources

Precipitation Data

- Daily gridded precipitation from nearest precipitation gauge
- CAPA dataset

•Discharge Data

- Water Survey of Canada (2016)
- Environment Canada gauging station

•Soil Moisture Data

 Weekly soil moisture anomaly data (Agriculture and Agrifood Canada) Complete SMOS and gauging station data series were only available for 2010-2014.

Only warm season data (May-November) were included.





Data Processing

 $S_n = \frac{\sum_g S_g A_{bg}}{\sum_g S_g}$

SMOS and **CAPA** Interpolation Methodology

Mean Soil Moisture (S) in catchment b is computed where Abg is fractional area of catchment b within grid cell g (Koster et al. 2000)

Semi-automated modified Chapman Baseflow separation algorithm

$$Q_{b}(i) = \frac{k}{2-k}Q_{b}(i-1) + \frac{1-k}{2-k}Q(i) \quad k = \left(\frac{Q}{Q_{5}}\right)^{1/t} \quad \text{Chapman, 1999}$$

 $Q_b(i)$ is baseflow at time *I* Q is the total stream flow

k is the baseflow recession parameter (a function of the slope of the correlation line [Q/Qo])

Calculating the Runoff Ratio for each watershed

Runoff Ratio = <u>Quickflow Discharge (m³/week)</u>

Precipitation(m/week) * Area (m²)



Runoff Ratio Over Catchments



Strength of Antecedent SMOS Soil Moisture-Runoff Relationship



- -0.20 0.01
- 0.01 0.18
- 0.18 0.38

43% of basins showed a statistically significant Rs value between the 0.38 - 0.59 I-week antecedent soil moisture anomaly and the weekly runoff ratio



2.000

Kilometers

1.000

500

0

Basin Characteristics

 R_s^2 Values for basin characteristics, runoff ratio and soil moisture-runoff correlations (N=76)

Average Annual

Soil Moisture-Runoff

Correlations calculated using 43 hydroclimatic, topographic, soil and landcover basin characteristics

		Correlation
Latitude	-0.09	-0.30
Longitude	0.05	0.43
Gross Drainage Area Km2	-0.06	0.21
Pfafstetter level	0.03	-0.16
Silt %	-0.45	0.12
Sand %	0.20	0.19
Clay %	-0.46	0.18
Gravel %	-0.07	-0.09
Organic Carbon %	-0.07	0.28
Drainage Density	-0.11	0.19
Average Annual Days of Precip	0.70	-0.01
Average Warm Season Precip (mm)	0.52	0.22
Precipitation Seasonality Index	-0.38	0.07
Minimum Slope	0.38	-0.36
Maximum Slope	0.62	-0.41
Mean Slope	0.62	-0.42
Compactness Ratio	0.17	-0.35
Min Elevation (m)	-0.10	-0.29
Max Elevation (m)	0.21	-0.41
Mean Elevation (m)	0.11	-0.34
%Settlement	-0.12	0.11
%Roads	-0.18	0.13
%Water	0.00	0.14
%Forest	0.43	-0.11
%Trees	0.14	0.02
%Treed Wetland	0.03	0.13
%Cropland	-0.26	0.05
%Grassland	-0.36	0.04
%Natural Grassland	0.49	-0.05
%Wetland	0.02	0.20
%Wetland Shrub	0.08	0.05
%Wetland Grass	0.02	-0.43
%Other (Rock, Ice, Barren)	0.45	-0.20
April Snowcover %	0.37	-0.20
May Snowcover %	0.48	-0.17
June Snowcover %	0.26	-0.16

Significant at the 95% level Significant at the 90% level

Key Findings

- ~40% of basins showed a significant correlation between the SMOS soil moisture anomaly and the basin runoff ratio
- Considerable variety in strength of relationship over individual basins
- Strong spatial patterns; relationship weakens in northern latitudes, mountainous and coastal regions
- Basin topography (slope, elevation) show the strongest controls on the relationship
- Landcover does not seem be have a dominant factor
- Method shows some promise as large scale tool to characterize controls on runoff across remote Canadian watersheds

Further analysis:

- Developing a predictive model for areas of high SMOS SM correlation
- Validation with supplemental basins



Thank You





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